







Using effective density spectrum to constrain crustal density profile of Vesta and Ceres

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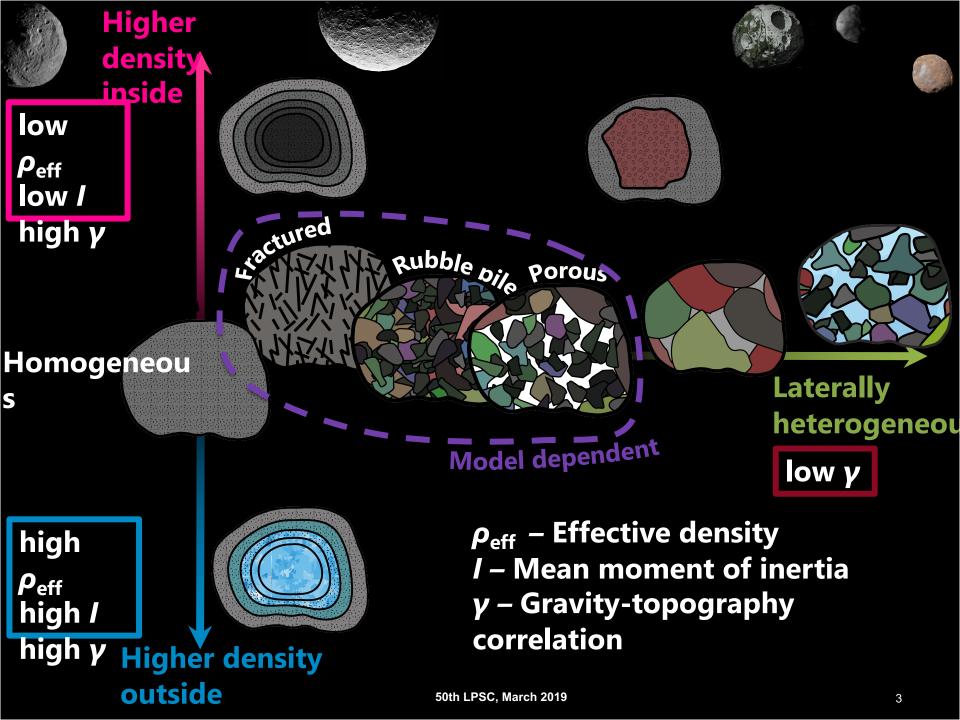








- **Gravity and topography in spherical harmonics**
- Derivation of effective density spectrum
- **Effective spectrum of Vesta**
- Why I thought it would work for Ceres
- and why it does not.











Spherical Harmonics

Shape

$$r(\phi, \lambda) = R_0 \sum_{n=0}^{\infty} \sum_{m=-n}^{n} A_{nm} Y_{nm} (\phi, \lambda)$$

Gravitational potential

$$U(r,\phi,\lambda) = \frac{GM}{r} \sum_{n=0}^{\infty} \sum_{m=-n}^{n} \left(\frac{R_0}{r}\right)^n C_{nm} Y_{nm}(\phi,\lambda)$$

U – gravitational potential

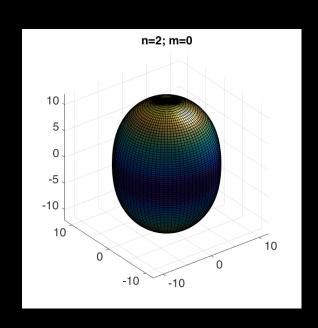
φ – latitude

 λ – longitude

r – radial distance

n – degree

m – order









Spherical Harmonics

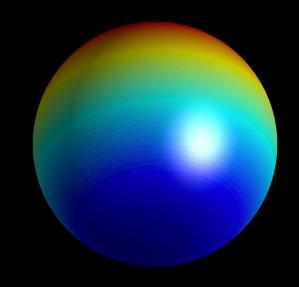
$$r(\phi, \lambda) = R_0 \sum_{n=0}^{\infty} \sum_{m=-n}^{n} A_{nm} Y_{nm} (\phi, \lambda)$$

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> RMS spectrum

$$M_n^{gg} = \sqrt{\frac{\sum_{m=-n}^n C_{nm}^2}{2n+1}}$$

$$M_n^{tt} = \sqrt{\frac{\sum_{m=-n}^{n} A_{nm}^2}{2n+1}}$$











Dawn geophysical data

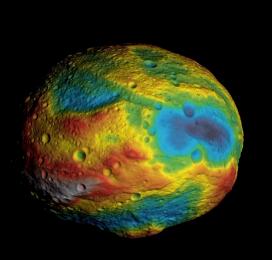
- Shape model
 - Stereophotogrammetry (SPG) from DLR
 - Stereophotoclinometry (SPC) from JPL
 - Mutually consistent with the accuracy much better than the spatial resolution of gravity field
- **Gravity field**
 - Accurate up to $n = 18 (\lambda = 93 \text{ km})$ for Vesta (Konopliv et al., 2014)
 - Globally accurate up to $n = 18 (\lambda = 174 \text{ km})$ for Ceres (Konopliv et al., 2017)
 - Locally accurate up to $n = 59 (\lambda = 50 \text{ km})$ (Park et al., in preparation)

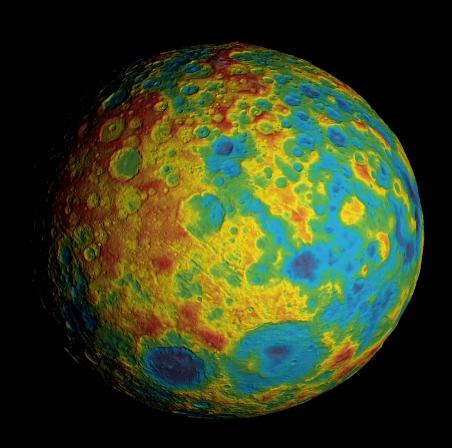
















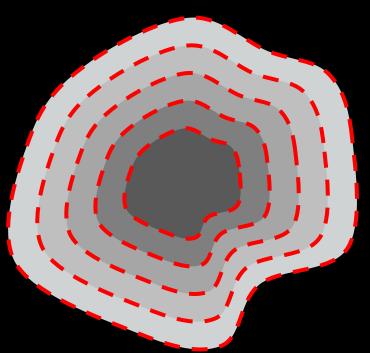


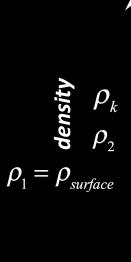


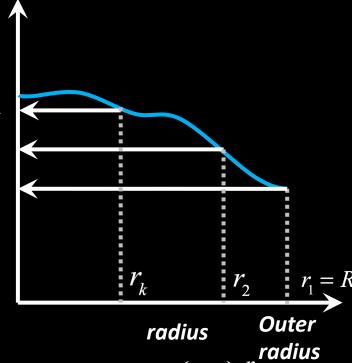
Effective density

$$\sigma_{nm}^{const} = \{ \overline{C}_{nm}^{const}, \overline{S}_{nm}^{const} \}$$

Gravity spherical harmonic coefficients referenced to the volumetric radius of a layer







Contribution of *k*-th layer to total gravity:

$$\sigma_{nm}^{const} \cdot \frac{4}{3} \pi r_k^3 (\rho_k - \rho_{k-1}) \cdot \frac{1}{M} \cdot \left(\frac{r_k}{R}\right)^n$$

Fractional mass of a layer 50th LPSC, March 2019

Upward continuation factor



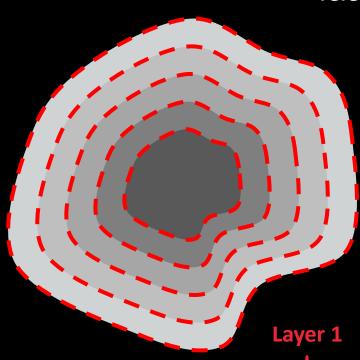


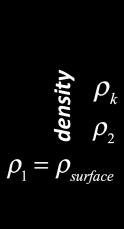


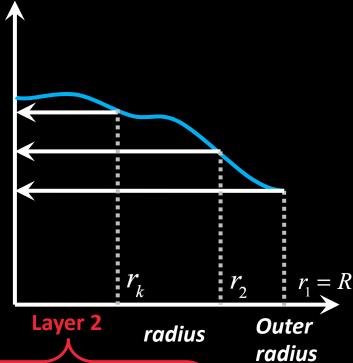
Effective density

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Gravity spherical harmonic coefficients referenced to the volumetric radius of a layer







$$\sigma_{nm} = \frac{1}{M} \sigma_{nm}^{const} \frac{4}{3} \pi r_1^3 \rho_1 \left(\frac{r_1}{R}\right)^n + \frac{1}{M} \sigma_{nm}^{const} \frac{4}{3} \pi r_2^3 (\rho_2 - \rho_1) \left(\frac{r_2}{Ronst}\right)^n + \dots$$

$$\sigma_{nm} = \frac{\sigma_{nm}}{\overline{\rho} R^{n+3}} \int_{R^+}^{0} r^{n+3} \frac{d\rho(r)}{dr} dr$$

$$\text{Layer k}$$

$$\sigma_{nm} = \frac{\sigma_{nm}^{Ronst}}{\overline{\rho}R^{n+3}} \int_{R^{+}}^{\tau} r^{n+3} \frac{d\rho(r)}{dr} dr$$
Layer k



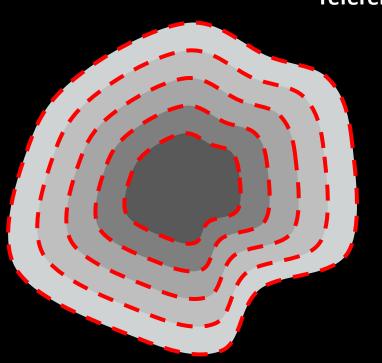


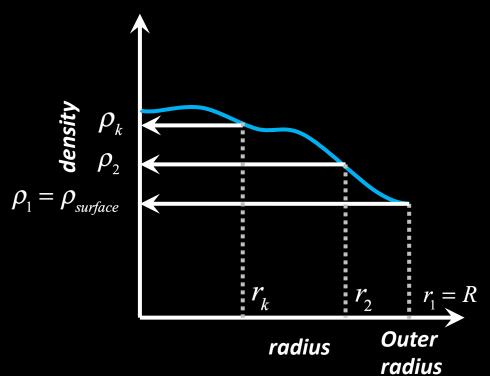




$$\sigma_{nm}^{const} = \{\overline{C}_{nm}^{const}, \overline{S}_{nm}^{const}\}$$

Gravity spherical harmonic coefficients referenced to the volumetric radius of a layer





Def: effective density

$$\tilde{\rho}_{n} = \frac{\sigma_{nm}}{\sigma_{nm}^{const}} \bar{\rho}$$

$$\tilde{\rho}_{n} = \rho_{surfage} \pm \frac{\sigma_{nm}^{const}}{R^{n}R^{n+3}} \int_{R^{-}R^{+}}^{0} r^{n+3} \frac{dap(r)}{dtr} dtr$$









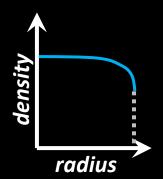
Simple density profiles

$$\tilde{\rho}_n = \rho_{surface} + \frac{1}{R^{n+3}} \int_{R^-}^0 r^{n+3} \frac{d\rho(r)}{dr} dr$$

$$\tilde{\rho}_{n} = \rho_{upper} + \left(\frac{r_{lower}}{r_{upper}}\right)^{n+3} \left(\rho_{upper} - \rho_{lower}\right)$$

$$\tilde{\rho}_n = \frac{n\rho_{surface} + 4\overline{\rho}}{n+4}$$

$$\tilde{\rho}_n = \frac{4n\overline{\rho} + 12\overline{\rho} - n\rho_{center}}{3(n+4)}$$



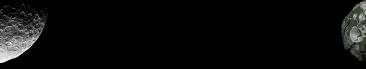
radius

radius

density

$$\tilde{\rho}_n = \rho_{surface} + (-1)^n \left(\frac{r}{R}\right)^{n+3} \Delta \rho e^{-R/d} \left[\Gamma(n+4) - \Gamma(n+4, -R/d)\right]$$







$$\tilde{\rho}_{n} = \rho_{surface} + \frac{1}{R^{n+3}} \int_{R^{-}}^{0} r^{n+3} \frac{d\rho(r)}{dr} dr$$

$$\tilde{\rho}_{_{0}}=\bar{\rho}$$

$$ilde{
ho}_{\scriptscriptstyle \infty} =
ho_{\scriptscriptstyle surface}$$









$$\lambda = \frac{\|g\|_R^2}{\|g\|_{\Omega}^2} = \frac{\int_R g^2 d\Omega}{\int_{\Omega} g^2 d\Omega} = \max$$

Spatial localization

$$\lambda = \frac{\|g\|_{L}^{2}}{\|g\|_{\infty}^{2}} = \frac{\sum_{l=0}^{L} \sum_{m=-l}^{l} h_{lm}^{2}}{\sum_{l=0}^{\infty} \sum_{m=-l}^{l} h_{lm}^{2}} = \max$$

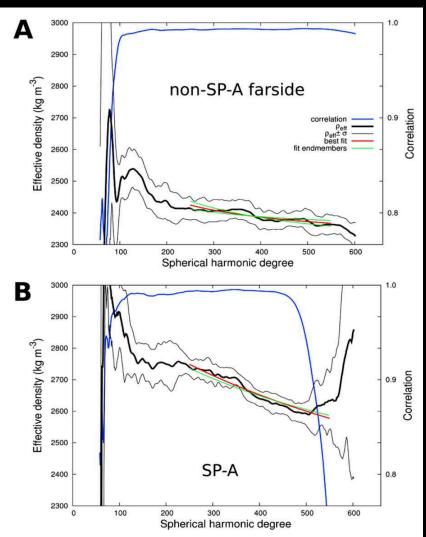
Spectral localization

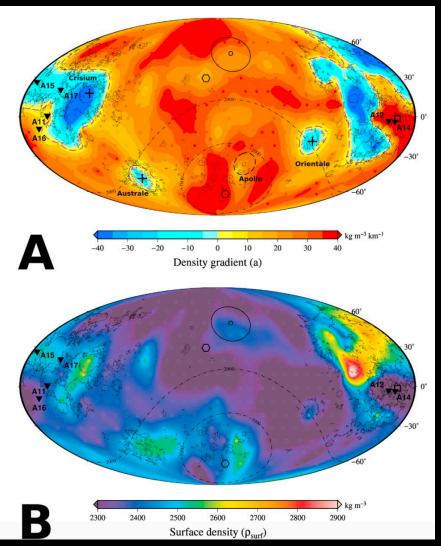
R – region of localization

- λ eigenvalue = concentration factor
- Simons et al., 2006 presents a way to find the eigenfunctions and shows that they are identical
- Allows localized estimates of crustal density



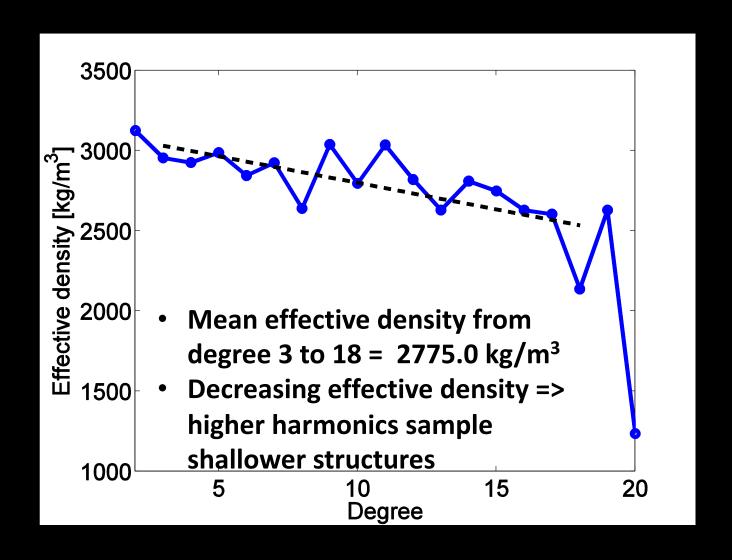






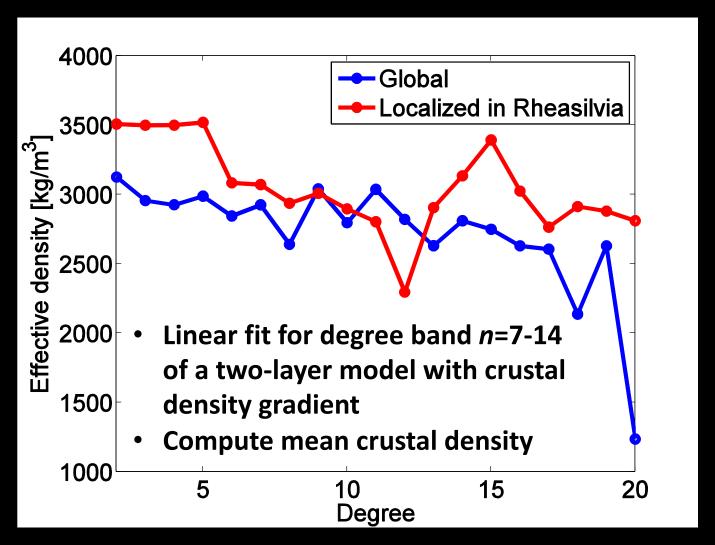






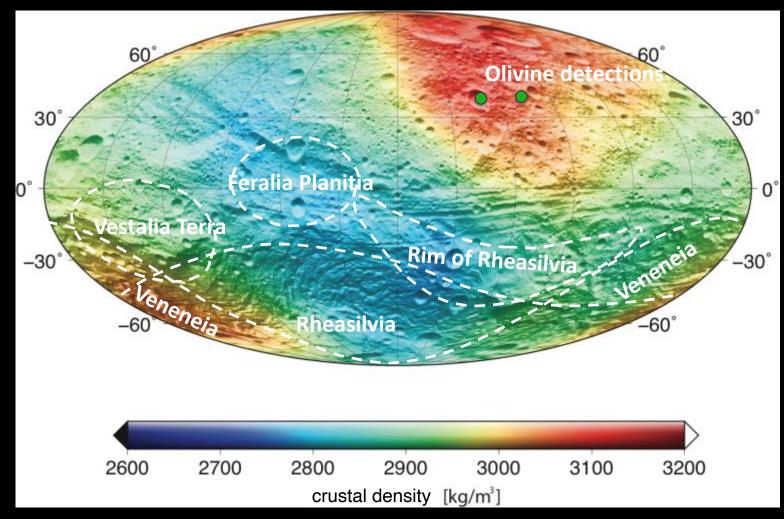






Interpolated crustal density

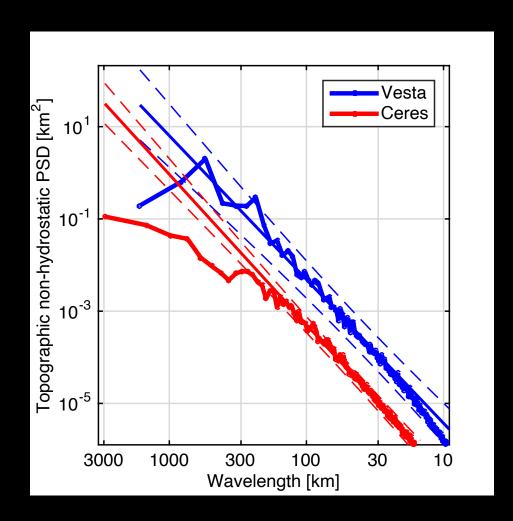
Bandwidth L = 5, window = 50° spherical cap





Evidence for viscous relaxation

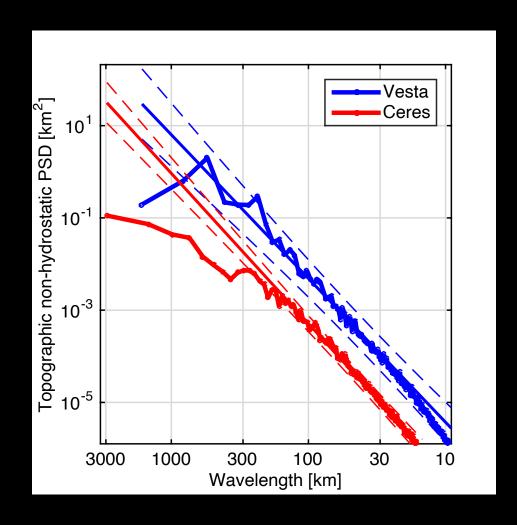
- More general approach: <u>study topography power</u> <u>spectrum</u>
- Power spectra for Vesta closely fits with the power law to the lowest degrees (λ < 750 km)
- Ceres power spectrum deviates from the power law at λ > 270 km







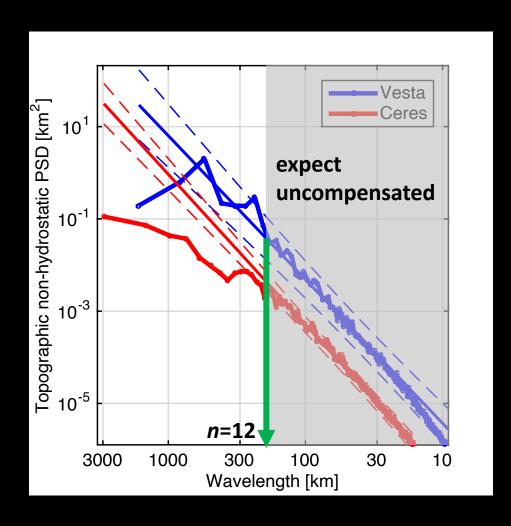
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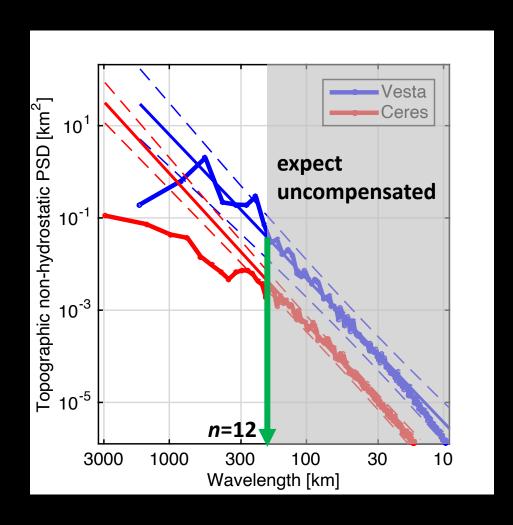
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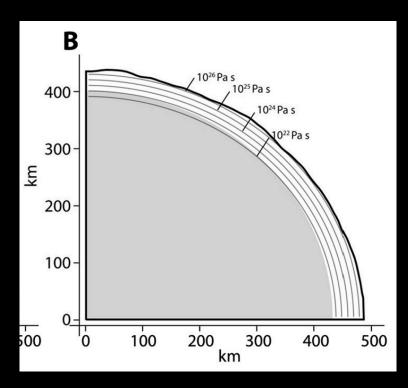
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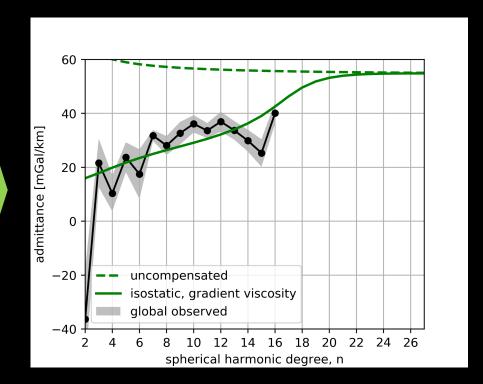
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Viscosity profile inferred by Fu et al., 2017 from topography relaxation



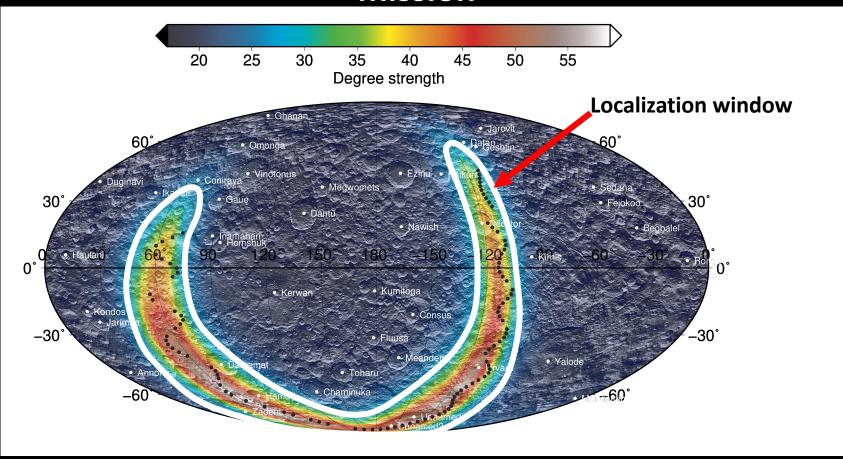






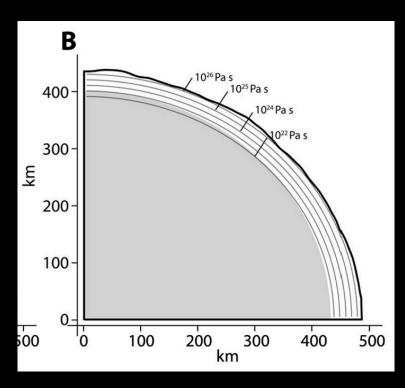


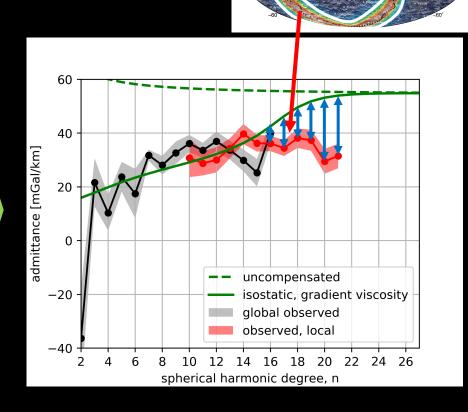






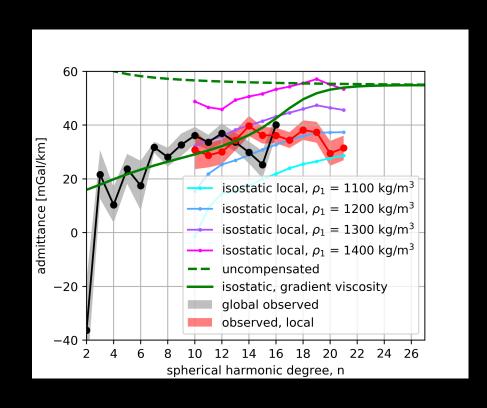
Viscosity profile inferred by Fu et al., 2017 from topography relaxation







- Localized admittance is consistent with isostatic compensation up to n = 21
- The mean crustal density is found to be 1233⁺⁴⁵₋₃₆ kg/m³ from the Dawn XM2 data.
- Possible signature of vertical density gradient at high degrees









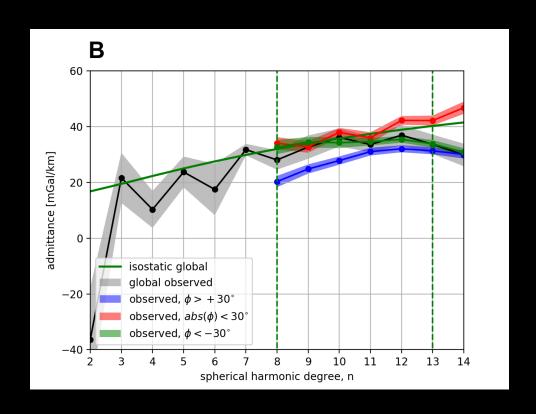




- Vesta uncompensated topography makes it possible to use effective density spectrum to constrain its density profile
 - Higher crustal density at the intersection of Rheasilvia and **Veneneia basins -> the region of deepest impact excavation**
 - Higher crustal density in the region of olivine detections
- Spectral-spatial localization of Ceres gravity and shape has not revealed the transition from compensated to uncompensated topography
 - This prevents using effective density spectrum for directly constraining crustal density of Ceres.
 - There remains a trade-off between compensation depth and crustal density.
 - Alternatively, a vertical density gradient in the crust of Ceres could be causing lower than uncompensated admittances



 Lower admittance near the poles, especially near the north pole



































$$\tilde{\rho}_n = \rho_{surface} + \frac{1}{R^{n+3}} \int_{R^-}^0 r^{n+3} \frac{d\rho(r)}{dr} dr$$

- **Assumptions:**
- mean density